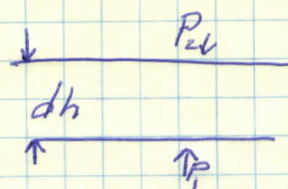


Garriot -

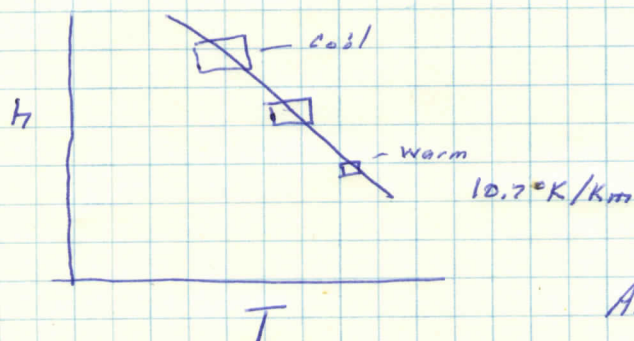


Air has ~ constant composition - to 100 KM

Temp $\propto H$

Temp variation $\bar{e}$ alt.

Adiabatic lapse rate is limiting rate of temp.



$\left. \begin{array}{l} 2^\circ \text{K}/1000' \\ \sim 6.5^\circ \text{K}/\text{KM} \end{array} \right\} \text{actual tropospheric rate}$

reduction in rate caused by presence of H_2O vapor

There is isothermal height region under ozone region
Ozone (filter) absorbs U.V. and is warmed

$$dp = dM \text{ gravity} = -\rho dh g$$

$$\rho = \text{number moles}$$

$$p = n k T$$

$$\frac{dp}{dh} = -n m g \left(\frac{kT}{kT} \right) \quad \boxed{H \equiv \frac{kT}{mg}}$$

$$\frac{dp}{dh} = -\frac{p}{H} \quad \ln p = -\frac{h}{H} + \text{const}$$

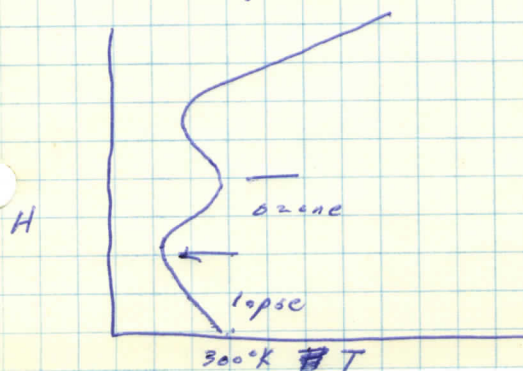
$$p = p_0 e^{-h/H} \quad H \equiv \text{scale height}$$

$$\text{now } z = \int_0^h \frac{dh}{H} \quad (\text{reduced height})$$

$$p = p_0 e^{-z}$$

$$n = n_0 \left(\frac{T_0}{T} \right) e^{-z}$$

Atmosphere is unstable if lapse rate is exceeded-

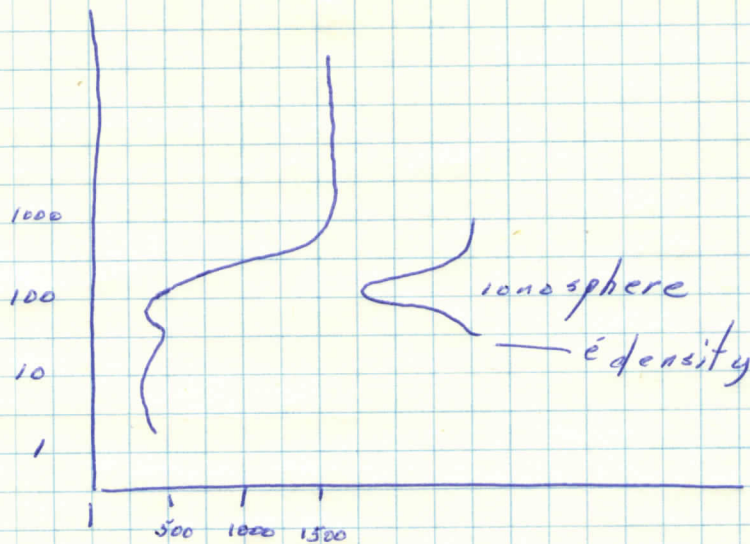


P-2

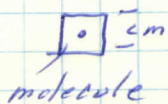
Aeronomy

Nov 22, '67

Once temp. is known other factors may be calculated to 100 KM - above this diffusion separation occurs $\bar{c} \uparrow$ He, etc.



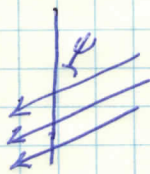
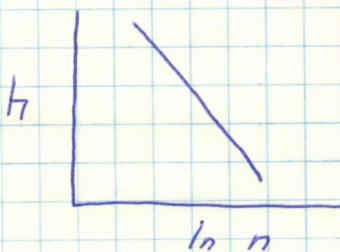
Compute altitude of dissociation
Optical X sect.



Flux - 1 photon/cm²

probability of absorption has cm² dimensions

May be considered in terms of area-length



$$dI \propto I_0 n ds \sigma$$

length
cross sect

$$\int_h^\infty \frac{dI}{I} = \int_h^\infty n \sigma \sec \psi dh$$

$$\ln I_0 - \ln I_h = \sigma \sec \psi \int_h^\infty n dh$$

$$I(h) = I_0 \exp. \left[\underbrace{-\sigma \sec \psi n H}_{-\tau} \right] \quad \int n dh = \int n_0 \left(\frac{T_0}{T} \right) e^{-z/h} \times \frac{Kt/mg}{Kt/mg}$$

$$= n_0 H_0$$

$$\sigma_{1,2} = 5 \times 10^{-19} \text{ cm}^2 \quad (1200 - 1800 \text{ \AA})$$

P-3

Acronomy

22 Nov.

$$\sigma \sim 5 \times 10^{-9} \text{ cm}^2 \text{ (120-180 mÅ)}$$

$$H \sim 10 \text{ km}, 10^6 \text{ cm}$$

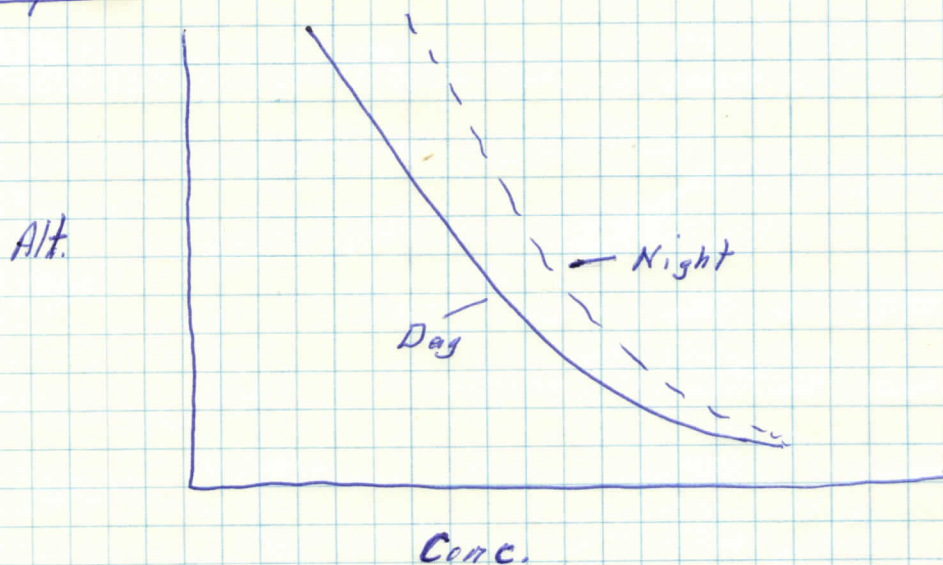
$$n \sim 2 \times 10^{12}$$

$$\text{Life of } O_2 \text{ molecule} \approx \frac{1}{I_0 \sigma}$$

$$\approx 3 \text{ days}$$

Undissociated O_2 must diffuse up

Temp. of thermosphere is dependent upon U.V. radiation

Ionosphere

Calc. Approx. Isothermal, only 1 constituent / ^{ionization} ~~x sect. not absorption~~ _{& prev. used}

$$q = \text{prod. rate of electrons + ions} = n \sigma I \text{ (elec/cm}^2\text{/sec)}$$

$$= n \sigma I_0 e^{-\tau}$$

ioniz. thresh. $O - 911 \text{ Å}^\circ$

$N_2 - 796 \text{ Å}^\circ$

$O_2 - 1025 \text{ Å}^\circ$

$$\frac{dq}{dh} = 0 \text{ Alt. of peak prod.}$$

$$= e^{-\tau} \frac{dn}{dh} n e^{-\tau} \left(-\frac{d\tau}{dh} \right)$$

$$\tau = \sigma n H \text{ sec } 4$$

Calc. cont.

$$* \quad l = n \sigma H \sec \psi = \tau$$

$$q = q_0 = n_0 \sigma \frac{I_\infty}{\infty}$$

$$q = \sigma I_\infty n_0 e^{-z}$$

Alt. at which
there is unit opt-
depth -

$$= \frac{I_\infty}{e H_0} \quad \text{value of peak prod. rate}$$

$$= q_0 \exp [1 - z - \sec \psi e^{-z}] \leftarrow \text{Chapman Prod. Funct.}$$

$$\sigma_{\text{ioniz}} = 10^{-15} \text{ cm}^2$$

These calculations assume plane parallel atmosphere -

Loss functions



$$\text{Loss Rate} = \alpha N^2$$

$$\frac{dN}{dt} = q - \alpha N^2 = 0 \quad N = \sqrt{\frac{q}{\alpha}} = \sqrt{\frac{q_0}{\alpha}} \exp \frac{1}{2} \left[1 - \frac{z}{2} - \sec \psi e^{-z} \right]$$

$$I_\infty = 10^{10} \text{ photons/cm}^2$$

$$N_{\max} \approx \sqrt{\frac{1000}{10^{-7}}} = 10^5$$

Compare Simplified Calc. to Meas. Ionosphere -